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Marine Chemistry

Volume 48, Issue 2, January 1995, Pages 157-182

Research paper

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Iron chemistry in seawater and its relationship to phytoplankton: a workshop report

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Received 26 July 1994; revision accepted 6 October 1994

1. Introduction

The ongoing debate about iron limitation of phytoplankton in the open ocean has highlighted how little we know about the marine chemistry of iron and its relationship to iron uptake by microorganisms. The idea that iron may be limiting phytoplankton was suggested and investigated by Gran (1931) but the topic suffered from analytical problems for almost sixty years (de Baar, 1994). The concept is now treated more seriously since the advent of trace metal clean procedures which permit the accurate measurement of dissolved and particulate iron (Bruland et al., 1979; Gordon et al., 1982; Landing and Bruland, 1987). These measurements demonstrate that in some ocean habitats dissolved iron concentrations in surface waters are as low as 20–30 pM (Martin et al., 1991), concentrations that are unlikely to support high phytoplankton biomass. The controversy over the role of iron in ocean productivity has stimulated the development of new methodologies for rapidly analyzing low-level iron concentrations in seawater (Elrod et al., 1991; O'Sullivan et al., 1991; Yokoi and van den Berg, 1992; Obata et al., 1993). Unfortunately, most of these novel methods provide little information about the ambient speciation of iron and hence its accessibility to phytoplankton. Critical assessment of the inter-

active relationship between iron and marine phytoplankton assemblages is hindered greatly by our current analytical resources.

Most of the iron debate centers on the remote HNLC regions of the subarctic Pacific, the equatorial Pacific and the Southern Ocean. These regions are characterized by a persistence of excess major nutrients (N, P) and low biomass relative to coastal systems having similar major nutrient concentrations. Because the root cause(s) for this situation may have important ramifications to the ocean–atmosphere exchange of CO₂ and global climate cycles (Martin, 1990), a special symposium was held (February 1991) to address what controls phytoplankton production in nutrient-rich areas of the open sea (see *Limnology and Oceanography*, Vol. 36). John Martin and his colleagues argued that an inadequate iron supply was the major factor (Martin et al., 1991). Results from bottle enrichment studies of theirs and others support this hypothesis (Martin and Fitzwater, 1988; de Baar et al., 1990; Martin et al., 1990a, 1994; Helbling et al., 1991; Greene et al., 1994; Kolber et al., 1994; Price et al., 1994), though alternate interpretations have suggested that other factors also play an important role (Dugdale and Wilkerson, 1990; Banse, 1991; Buma et al., 1991; Mitchell et al., 1991; Nelson and Smith, 1991) [see the reviews of Cullen (1991) and de Baar (1994)]. The recent

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